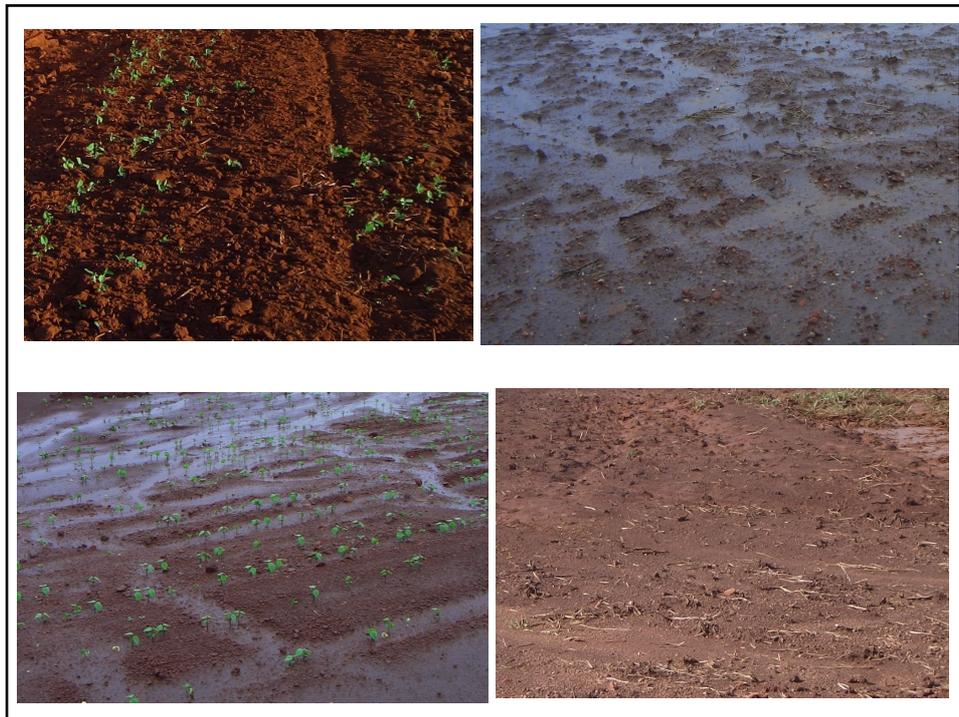


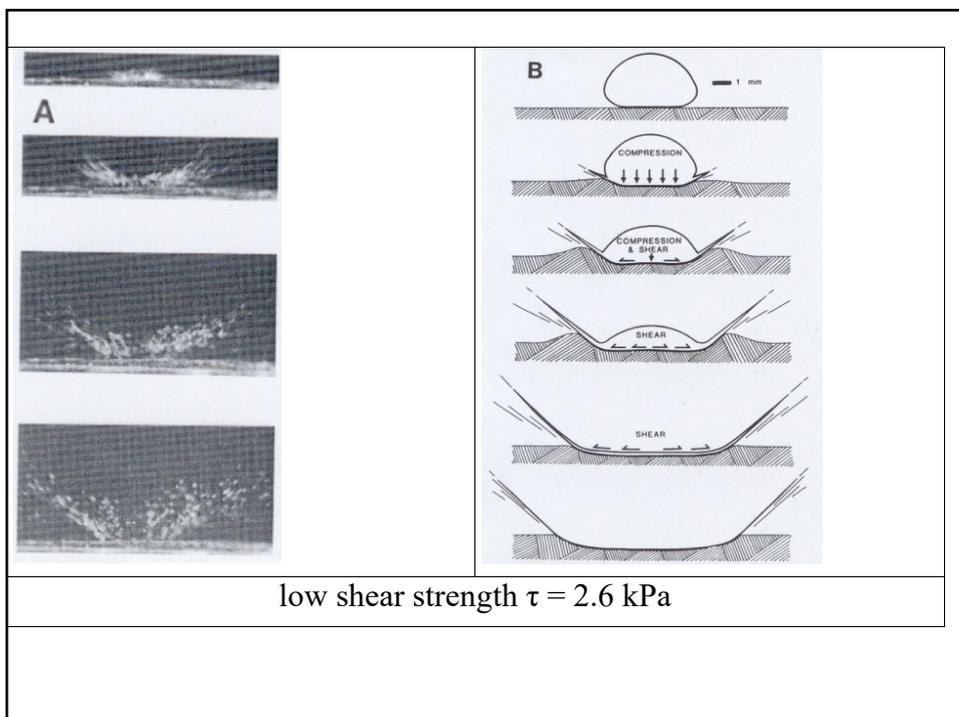
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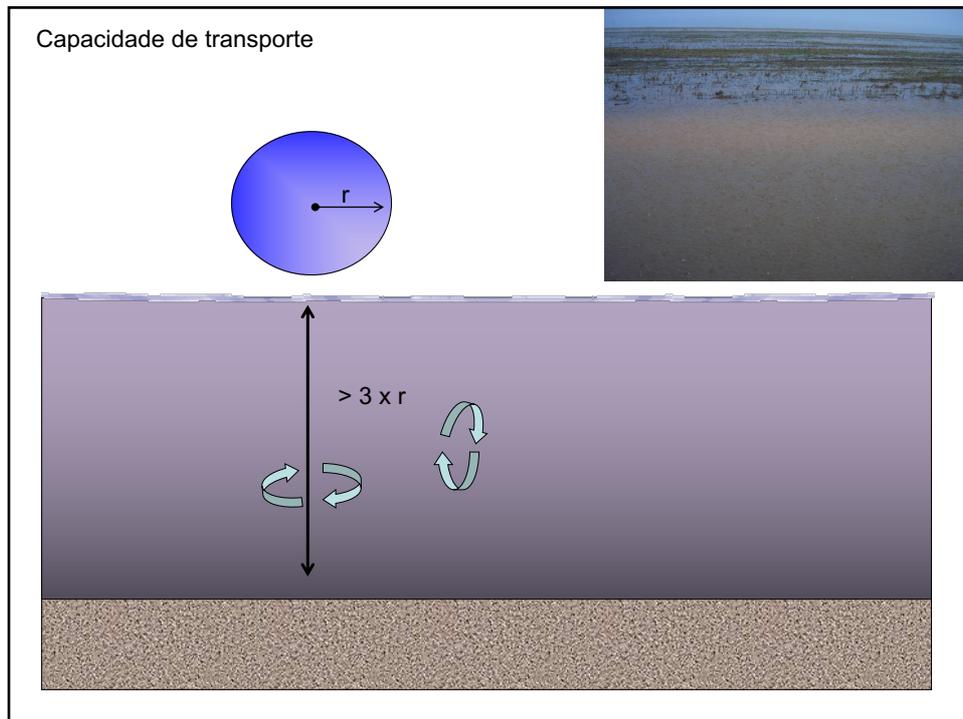
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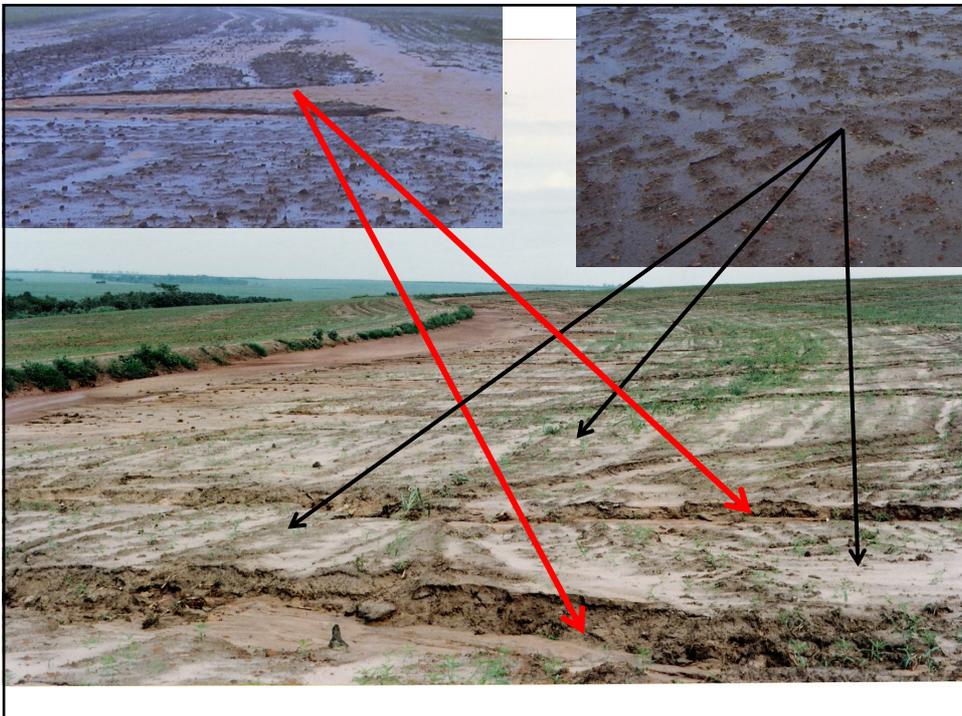
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7



8

Erosividade da Chuva

Quantidade de Chuva

- Dados de chuva em totais ou médias mensais ou anuais pouco significam em relação à erosão.
- Dados de chuva totais diários, limitados a observações a cada 24hs, pouco significado em relação à erosão.

Distribuição da chuva

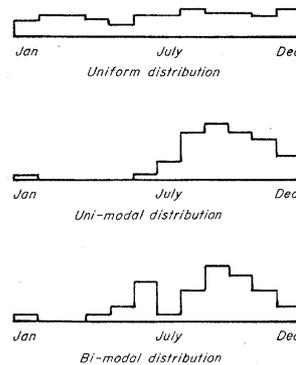


FIGURE 3.1 Patterns of distribution of annual rainfall

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Intensidade da Chuva:

- fator mais importante
- quanto maior a intensidade da chuva maior a perda por erosão

exemplo: para uma chuva total de 21mm, uma intensidade de 7,9mm/h produziu uma perda de terra 100 vezes maior que uma de 1mm/h

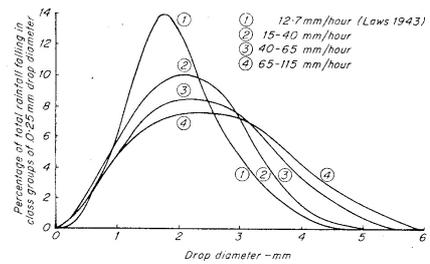


FIGURE 3.4 Drop size distribution at low and medium intensities

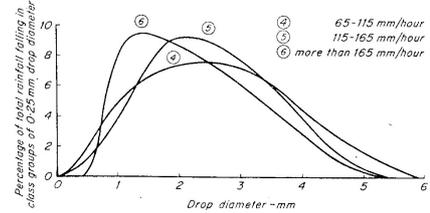


FIGURE 3.5 Drop size distribution at high intensities

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Table 1
 Reported relationships between time-specific kinetic energy (KE_{time}) and intensities (I) of rains. The expression of KE_{time} not reported in this table can easily be obtained by the relation (see Eq. (3)) $KE_{time} (J m^{-2} min^{-1}) = KE_{max} (J m^{-2} h^{-1})/I (DSD, \text{drop-size distribution; n.a., not available})$

Reference	$KE_{time} (J m^{-2} h^{-1}) - I (mm h^{-1})$ relation	Location	Range of $I (mm h^{-1})$
Bollinne et al., 1984	$12.32I + 0.56I^2$	Belgium	0.27–38.6
Brandt, 1990	$I (8.95 + 8.44 \log_{10} I)$	USA, DSD from Marshall and Palmer (1948)	n.a.
Brown and Foster, 1987	$29I (1 - 0.72e^{-0.05I})$	USA	0–250
Carter et al., 1974	$11.32I + 0.5546I^2 - 0.5009 \times 10^{-3}I^3 + 0.126 \times 10^{-4}I^4$	South Central USA	1–250
Cerro et al., 1998	$38.4I (1 - 0.538e^{-0.05I})$	Barcelona, Spain	n.a.
Coimbra and Tomás, 1995	$35.9I (1 - 0.559e^{-0.05I})$	Portugal	0–120
Hudson, 1965	$29.86I (-4.29)$	Zimbabwe	n.a.
Jayawardena and Rezaur, 2000b	$36.8I (1 - 0.691e^{-0.05I})$	Hong Kong	0–150
Kinnell, 1981	$I(17.12I + 5.229 \log_{10} I)$	Miami, Florida	1.89–309
	$30.132I (-5.484)$		
	$29.51I (1 - 0.281e^{-0.018I})$		
Kinnell, 1981	$I (9.705 + 9.258 \log_{10} I)$	Rhodesia (from Hudson (1961))	18.5–228.6
	$29.863I (-4.287)$		
	$29.22I (1 - 0.894 e^{-0.007I})$		
McGregor and Mutchler, 1976	$I (27.3 + 21.68e^{-0.004I} - 41.26e^{-0.002I})$	Mississippi, USA	n.a.
Onaga et al., 1988	$I (9.81 + 10.6 \log_{10} I)$	Okinawa, Japan	n.a.
Park et al., 1980	$21.1069I^{0.705}$	USA	n.a.
Renard et al., 1992	$29I (1 - 0.72e^{-0.05I})$	USA	n.a.
Rosewell, 1986	$29I (1 - 0.596e^{-0.006I})$	Gunnedah, Australia	1–145.9
Rosewell, 1986	$26.35I (1 - 0.669e^{-0.005I})$	Brisbane, Australia	1–161.2
Rosewell, 1986	$24.48 (I - 1.253)$	Melbourne, Australia	n.a.
Rosewell, 1986	$24.80 (I - 1.292)$	Cowra, Australia	n.a.
Sempere-Torres et al., 1992	$34I - 190$	Cévennes, France	20–100
Smith and De Veaux, 1992	$13I^{1.23}$	Oregon, USA	n.a.
Smith and De Veaux, 1992	$18I^{1.26}$	Alaska, USA	n.a.
Smith and De Veaux, 1992	$11I^{1.17}$	Arizona, USA	n.a.
Smith and De Veaux, 1992	$10I^{1.18}$	New Jersey, USA	n.a.
Smith and De Veaux, 1992	$11I^{1.14}$	North Carolina, USA	n.a.
Smith and De Veaux, 1992	$11I^{1.14}$	Florida, USA	n.a.
Steiner and Smith, 2000	$11I^{1.25}$	Northern Mississippi, USA	n.a.
Tracy et al., 1984	$210Ie^{-0.0056I} - 211.8$ if $I < 76 \text{ mm h}^{-1}$	Arizona	n.a.
Uijlenhoet and Stricker, 1999a	$7.20I^{0.32}$	Based on Marshall and Palmer parameterisation	n.a.
	$8.53I^{0.29}$		
	$8.46I^{0.17}$		
	$8.89I^{0.28}$		
	$10.8I^{0.96}$		
	$7.74I^{0.35}$		
Usón and Ramos, 2001	$23.4I - 18$	NE Spain	< 20
Wischmeier and Smith, 1958	$I (11.87 + 8.73 \log_{10} I)$	Washington, USA; DSD from Laws and Parsons (1943)	n.a.
Zanchi and Torri, 1980	$I (9.81 + 11.25 \log_{10} I)$	Italy	n.a.

Salles et al., 2002

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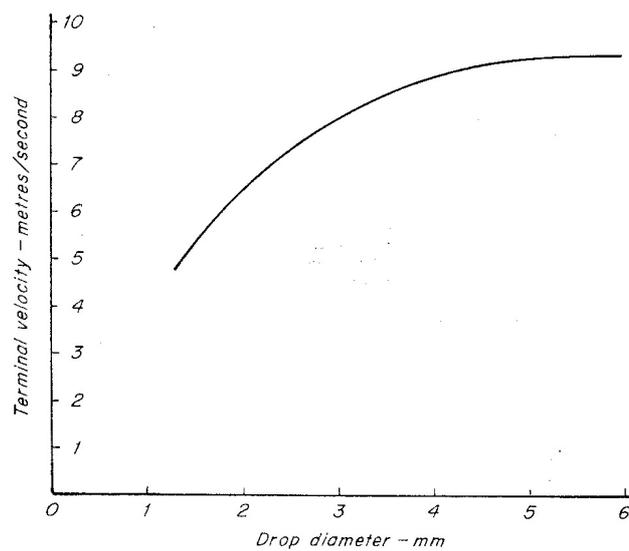


FIGURE 3.8 The terminal velocity of raindrops (data from LAWS 1941)

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Energia Cinética da Gota de Chuva

É muito importante conhecer a força com que a chuva impacta contra o solo. A energia cinética é uma função da massa e da velocidade:

$$E_c = \frac{1}{2}mv^2$$

Energia cinética da chuva e da enxurrada:

	Chuva	Enxurrada
Massa	Suponha uma massa de queda da chuva = R	Suponha 25% de enxurrada, e a massa da enxurrada = R/4
Velocidade	Suponha uma velocidade de 8m/s	Suponha a velocidade de escoamento na superfície de 1m/s
Energia Cinética	$1/2 \times R \times (8)^2 = 32R$	$1/2 \times R/4 \times (1)^2 = R/8$

Desta forma a chuva tem 256 vezes mais energia cinética que a enxurrada

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Duração da chuva:

- complemento da intensidade
- a combinação dos dois determina a chuva total
- importante para determinar o limite no qual o solo para de infiltrar a água e começa a enxurrada.

Frequência de chuva:

- fator importante que influi nas perdas de terra por erosão.

Se intervalos são curtos \Rightarrow o teor de umidade do solo é alta \Rightarrow enxurradas mais volumosas (inclusive com chuvas de menor intensidade)

Se intervalos maiores \Rightarrow o solo esta seco \Rightarrow boa infiltração inicial \Rightarrow pouca enxurrada com chuvas de baixa intensidade.

Períodos de estiagem \Rightarrow diminuição da proteção vegetal por falta de umidade \Rightarrow maiores riscos de erosão nas primeiras chuvas

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Modelos estatísticos:

Zingg, 1940	$A = C' L^{0.6} S^{1.4}$
Smith, 1941	$A = C'' L^{0.6} S^{1.4} P$
Browning, 1947	$A = C''' L^{0.6} S^{1.4} P$
Musgrave, 1947	$A' = (P_{30}/1.25)^{1.75} K' (L/72.6)^{0.35} (S/10)^{1.35} C^*$
USLE, 1965	$A = EI_{30} K (L/72.6)^{0.5} (0.065 + .045 S + .0065 S^2) C P$
USLE, 1978	$A = EI_{30} K (L/72.6)^{0.5} (65.4 \sin^2 \Theta + 4.56 \sin \Theta + 0.065) C P$
RUSLE, 1997	$A = EI_{30} K (L/72.6)^m (a \sin \Theta + b) C P$

Energia cinética
Intensidade máxima

↓

Conversão estatística
utilizando funções do
volume e da
intensidade de
precipitação

A – Soil loss in tons/acre
 A' – Soil loss in inches/year
 C', C'', C''' – Coefficients
 C* – vegetal cover factor
 P₃₀ – Maximum Precipitation amount (inches) falling in 30 minutes in a storm
 K', K – Soil erodibility factors
 L – Slope length in feet
 S – Slope in percent
 Θ – Slope angle in degrees
 C – Cropping management factor
 E – Storm rainfall energy in hundreds of foot-tons per acre
 I₃₀ – Maximum rainfall intensity in a 30 minute period within a storm in inches per hour
 P – Conservation practice factor
 M – Exponent on length term-values depend on slope or slope and rill/interrill ratio
 a, b – coefficients in function making up slope term – values depend on slope

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USLE

A = RKLSCP

A = Long term average annual soil loss
 R = Rainfall erosivity
 K = Soil erodibility
 L and S = Slope factors
 C = Crop factor
 P = Management factor

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Rainfall Erosivity (R)

Wischmeier and Smith, 1978

$$e = 0.119 + 0.0873 \log_{10} i \tag{1}$$

where e is the unitary kinetic energy ($\text{MJ ha}^{-1} \text{mm}^{-1}$) and i represents the segments of rainfall intensity (mm h^{-1}).

$$EI_{30} = Ect I_{30} \tag{2}$$

where EI_{30} is the rainfall erosivity index ($\text{MJ mm ha}^{-1} \text{h}^{-1}$), Ect is the total kinetic energy of the rain (MJ ha^{-1}), and I_{30} is the maximum rain intensity during a 30-minute period (mm h^{-1}).

RUSLE

$$R = \frac{1}{n} \sum_{j=1}^n \sum_{k=1}^{m_j} (EI_{30})_k \tag{3}$$

where R is the average of the annual rainfall erosivity ($\text{MJ mm ha}^{-1} \text{h}^{-1} \text{yr}^{-1}$), n is the number of years of records, m_j is the number of erosive events in a given year j , and EI_{30} is the rainfall erosivity index of a single event k .

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CATENA

Rainfall erosivity map for Brazil
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 Received 5 February 2003; received in revised form 15 October 2003; accepted 24 November 2003

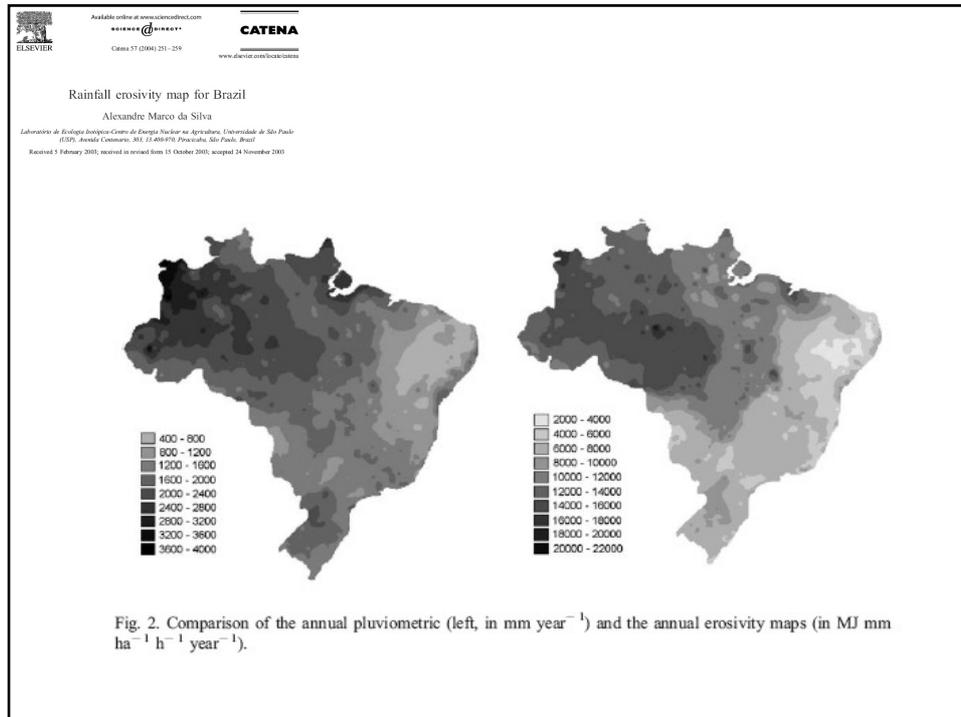


Number	Equation	Author(s)
1	$R_x = 3.76 * \left(\frac{M_x^2}{P}\right) + 42.77$	Oliveira Jr. and Medina (1990)
2	$R_x = 36.849 * \left(\frac{M_x^2}{P}\right)^{1.002}$	Morais et al. (1991)
3	$R_x = (0.66 * M_x) + 8.88$	Oliveira Jr. (1988)
4	$R_x = 42.307 * \left(\frac{M_x^2}{P}\right) + 69.763$	Silva (2001)
5	$R_x = 0.13 * (M_x^{2.24})$	Leprun (1981)
6	$R_x = 12.592 * \left(\frac{M_x^2}{P}\right)^{0.4026}$	Val et al. (1986)
7	$R_x = 68.73 * \left(\frac{M_x^2}{P}\right)^{0.841}$	Lombardi Neto and Moldenhauer (1992)
8	$R_x = 19.55 + (4.20 * M_x)$	Rufino et al. (1993)

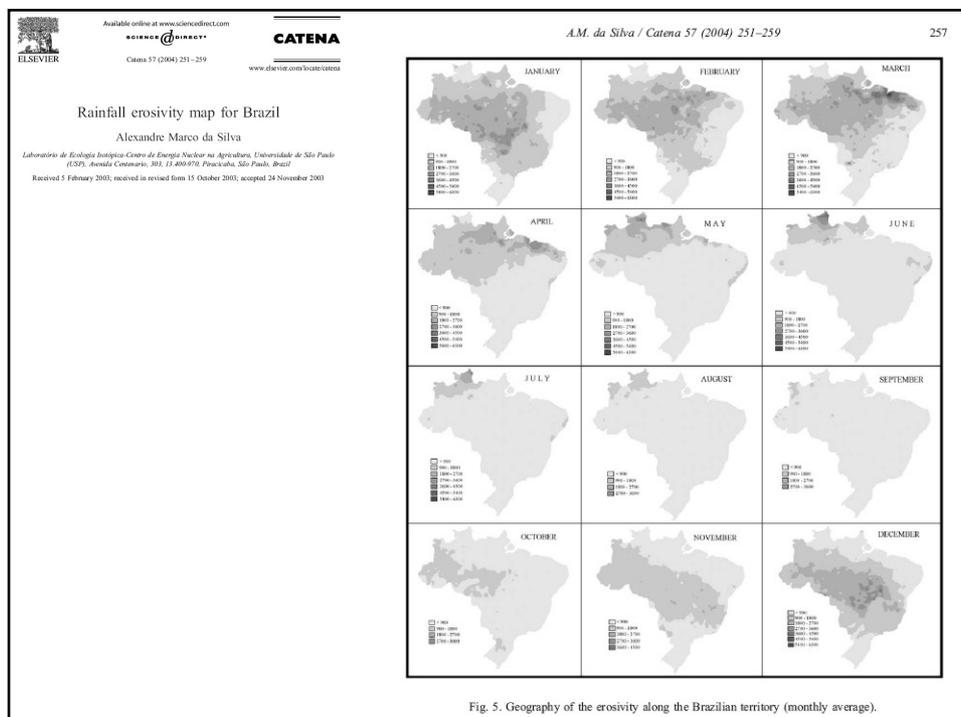
x = índice mensal
M = Precipitação média mensal
P = Precipitação média anual

Fig. 1. Equations used to determine the monthly/annual values of the erosivity according to the area of the territory and their respective authors. Eqs. number (1), (2), (4), (6), and (7) were based from the Fournier's model. Eqs. (3) and (8) are linear models and Eq. (5) is an exponential model. R_x is R factor ($\text{MJ mm ha}^{-1} \text{h}^{-1} \text{year}^{-1}$) for month x , M_x is average monthly precipitation depth (mm), and P is average annual precipitation (mm) (Silva, 2001).

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Equações da Erosividade no Brasil

Table 1 (continued)

Latitude	Longitude	City	States	Years	Altitude	Precipitation	R factor	Equations	R ²	Authors
22° 28' 48" S	42° 39' 36" W	Conc. Macabu	RJ	15	40	1915	7961	$E_{10} = 39.86 + 37.90 (MP)$	0.91	Gonçalves et al., 2006
22° 28' 48" S	43° 0' 0" W	Magé	RJ	16	640	3006	15,806	$E_{10} = 146.28 + 46.37 (MP)$	0.70	Gonçalves et al., 2006
22° 51' 0" S	42° 32' 00" W	Saquarema	RJ	15	10	1252	5448	$E_{10} = -13.36 + 50.02 (MR)$	0.65	Gonçalves et al., 2006
22° 55' 12" S	43° 25' 12" W	Rio de Janeiro	RJ	17	40	1260	4439	$E_{10} = 38.9 + 37.76 (MP)$	0.79	Gonçalves et al., 2006
22° 57' 36" S	43° 16' 48" W	Rio de Janeiro	RJ	16	460	2170	9331	$E_{10} = -76.27 + 53.31 (MR)$	0.40	Gonçalves et al., 2006
22° 42' 38" S	43° 52' 41" W	Pirai	RJ	18	462	-	6772	-	-	Machado et al., 2008
30° 23' 0" S	56° 26' 0" W	Quaraí	RS	38	100	1513	9292	$E_{10} = -47.35 + 82.72 (MR)$	0.84	Bazzano et al., 2007
32° 01' 0" S	52° 09' 0" W	Rio Grande	RS	23	15	1162	5135	Non-significant correlation	-	Bazzano et al., 2010
28° 39' 0" S	56° 0' 0" W	São Borja	RS	48	99	1540	9751	$E_{10} = 99.646 + 63.874 (MR)$	0.77	Cassol et al., 2008
								$E_{10} = 55.564 (MP)^{1.064}$	0.84	
30° 32' 0" S	52° 31' 0" W	Enc. do Sul	RS	31	420	1279	5534	Non-significant correlation	-	Eltz et al., 2011
29° 45' 0" S	57° 05' 0" W	Uruguaiana	RS	29	74	1171	8875	$E_{10} = -96.735 + 81.967 (MP)$	0.94	Hickmann et al., 2008
28° 33' 0" S	53° 54' 0" W	Ijuí	RS	31	448	1667	8825	$E_{10} = 330.86 + 34.54 (MP)$	0.40	Cassol et al., 2007
								$E_{10} = 168.65 (MP)^{0.75}$	0.53	
27° 51' 0" S	54° 29' 0" W	Santa Rosa	RS	29	273	1832	11,217	$E_{10} = 254.71 + 44.927 (MR)$	0.41	Mazurana et al., 2009
								$E_{10} = 118.52 (MP)^{0.624}$	0.50	
27° 24' 0" S	51° 12' 0" W	Campos Novos	SC	10	947	1754	6329	$E_{10} = 238.585 + 22.626 (MR)$	0.50	Bertol, 1994
								$E_{10} = 59.265 (MP)^{1.067}$	0.86	
27° 49' 0" S	50° 20' 0" W	Lages	SC	10	953	1549	5790	-	-	Bertol et al., 2002
22° 37' 0" S	52° 10' 0" W	Teófilo, Campinas	SP	19	255	1282	7172	$E_{10} = 106.8183 + 46.9562 (MP)$	0.93	Colodro et al., 2002
22° 31' 12" S	47° 2' 40" W	Campinas	SP	22	670	1280	6738	$E_{10} = 68.730 (MP)^{0.641}$	0.98	Lombardi Neto and Moldenhauer, 1992
23° 13' 0" S	49° 14' 0" W	Pirajua	SP	23	571	1482	7074	$E_{10} = 72.5488 (MR)^{0.668}$	0.93	Roque et al., 2001
24° 17' 0" S	47° 57' 0" W	Seze Barras	SP	9	30	1434	12,664	$E_{10} = 316.20 + 55.40 (MP)$	0.98	Silva et al., 2009b
24° 24' 0" S	47° 45' 0" W	Juquiá	SP	7	60	824	6145	$E_{10} = 207.21 + 40.65 (MP)$	0.90	Silva et al., 2009b
21° 16' 58" S	47° 0' 36" W	Mooca	SP	-	-	-	-	$E_{10} = 111.173 (MR)^{0.591}$	0.98	Carvalho et al., 1991

Years = length of record, altitude (m), P = average annual precipitation (mm), R = R factor (MJ mm ha⁻¹ h⁻¹ yr⁻¹), and (-) not available. States by region: north (Amazonas, AM and Pará, PA); northeast (Ceará, CE and Pernambuco, PE); central-west (Mato Grosso do Sul, MS; Mato Grosso, MT and Goiás, GO); southeast (Espírito Santo, ES; Minas Gerais, MG; Rio de Janeiro, RJ and São Paulo, SP) and south (Paraná, PR, Rio Grande do Sul, RS and Santa Catarina, SC).

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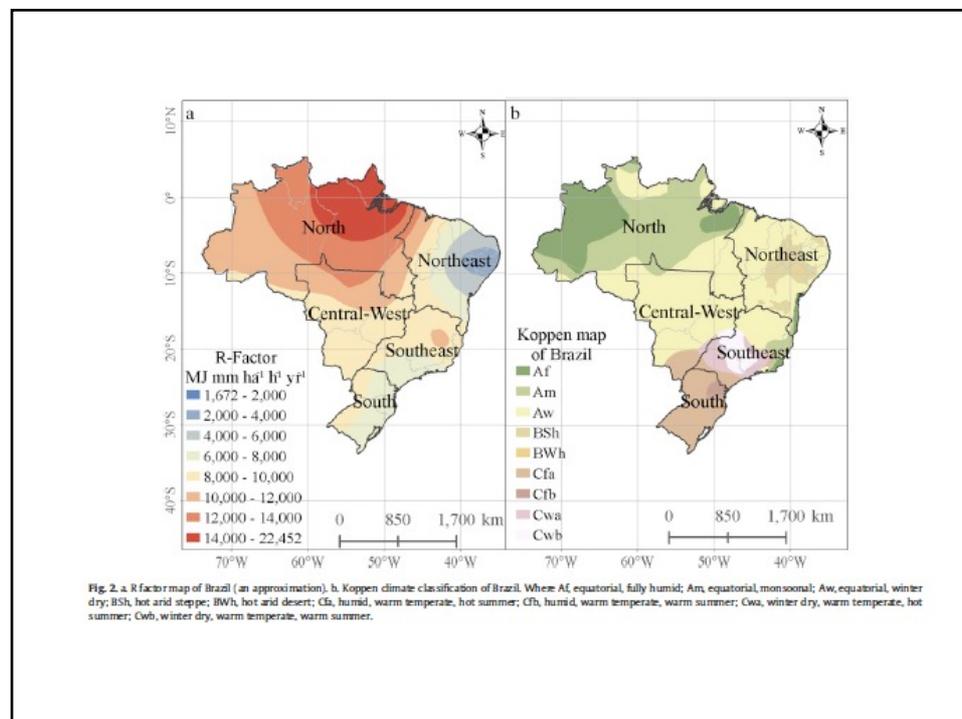


Fig. 2. a. R factor map of Brazil (an approximation). b. Koppen climate classification of Brazil. Where: Af, equatorial, fully humid; Am, equatorial, monsoonal; Aw, equatorial, winter dry; BSh, hot arid steppe; BWh, hot arid desert; Cfa, humid, warm temperate, hot summer; Cfb, humid, warm temperate, warm summer; Cwa, winter dry, warm temperate, hot summer; Cwb, winter dry, warm temperate, warm summer.

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Table 2
Range of rainfall erosivity values for several locations of the world.

Locate	Erosivity (MJ mm ha ⁻¹ h ⁻¹ yr ⁻¹)	Source
Tropical sites		
Honduras	2980-7297	Mikhailova et al. (1997)
Peninsular Malaysia	9000-14000	Shamshad et al. (2008)
Colombian Andes	10,409-15,975	Hoyos et al. (2005)
El Salvador Republic	7196-17856	Silva et al. (2011)
Southeastern Nigeria	12,814-18,611	Obi and Salako (1995)
Brazil	1672-22452	Present paper
Australia's tropics	1080-33500	Yu (1998)
Temperate sites		
Slovenia	1318-2995	Mikos et al. (2006)
Mediterranean region	100-3203	Diodato and Illecchi (2010)
Northeastern Spain	40-4900	Angulo-Martinez et al. (2009)
Switzerland	124-5611	Meusburger et al. (2011)
Korea	2109-6876	Lee and Hoo (2011)
Central Chile	50-7400	Bonilla and Vidal (2011)
United States	85-11,900	Renard and Heimund (1994)

Table 3
Classifications for the interpretation of the annual erosivity index of Brazil.

Erosivity ^a (MJ mm ha ⁻¹ h ⁻¹)	Erosivity class	Observed data (%)
R ≤ 2452	Low erosivity	26
2452 < R ≤ 4905	Medium erosivity	132
4905 < R ≤ 7357	Medium-strong erosivity	316
7357 < R ≤ 9810	Strong erosivity	237
R > 9810	Very strong erosivity	289

^a Source: Carvalho (2008), modified to S.I. metric units according to Foster et al. (1981).

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Process Based Models: Mass balance differential equation

$$\delta(cq)/\delta x + \delta(ch)/\delta t + S = 0$$

c = Sediment concentration (kg/m³)
q = Runoff discharge (m²/s)
x = Distance in the direction of flow (m)
h = Depth of flow (m)
t = Time (s)
S = Source/sink term (kg/(m²s))

WEPP

$$dG/dx = Dr + Di$$

G = Sediment load per unit width in the flow (kg/(m s)) = cq
Dr = net rill erosion rate per unit area of rill bottom (kg/(m² s)) (+ or -)
Di = Interrill sediment delivery to rill (kg/(m² s))

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Entrada de dados pelo WEPP

da	mo	year	prcp (mm)	dur (h)	tp	ip	tmax (C)	tmin (C)	rad (l/d)	w-vel (m/s)(Deg)	w-dir (C)	tdew
1	1	1	22.6	2.42	0.04	5.81	30.0	18.9	214.	5.5	234.	22.3
2	1	1	13.3	1.64	0.26	6.93	29.0	18.2	374.	5.9	188.	20.7
3	1	1	41.9	1.57	0.77	7.01	29.4	18.3	578.	5.0	86.	18.0
4	1	1	0.0	0.00	0.00	0.00	32.1	18.1	464.	5.9	264.	18.1
5	1	1	87.5	4.97	0.28	9.70	30.3	20.0	353.	5.4	179.	21.2
6	1	1	11.8	1.91	0.05	6.32	31.7	21.3	327.	6.1	183.	21.4
7	1	1	0.0	0.00	0.00	0.00	29.5	20.4	410.	4.8	45.	20.0
8	1	1	0.0	0.00	0.00	0.00	31.5	19.8	591.	3.8	97.	20.2
9	1	1	8.3	3.99	0.61	8.28	30.4	20.0	563.	6.4	357.	20.8
10	1	1	10.2	3.68	0.23	7.55	30.9	18.6	526.	4.2	305.	21.2
11	1	1	0.0	0.00	0.00	0.00	30.9	18.3	631.	0.0	0.	17.0
12	1	1	0.0	0.00	0.00	0.00	29.8	18.5	725.	4.0	324.	21.2
13	1	1	0.0	0.00	0.00	0.00	29.3	19.8	474.	7.1	273.	22.4

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Efeito de chuvas extremas:

Edwards and Owness (1991):

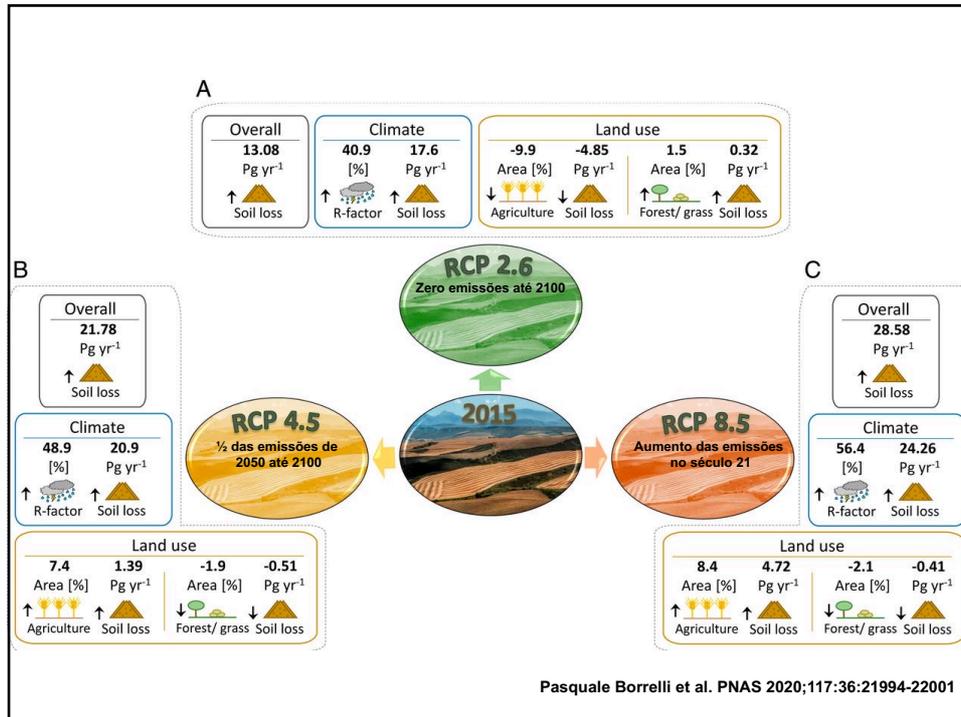
“Com mais de 4.000 eventos chuvosos monitorados durante o período do estudo (28 anos), os cinco (5) maiores eventos erosivos em cada microbacia (229) foram responsáveis por 66% da erosão total.”

Ranking do Evento

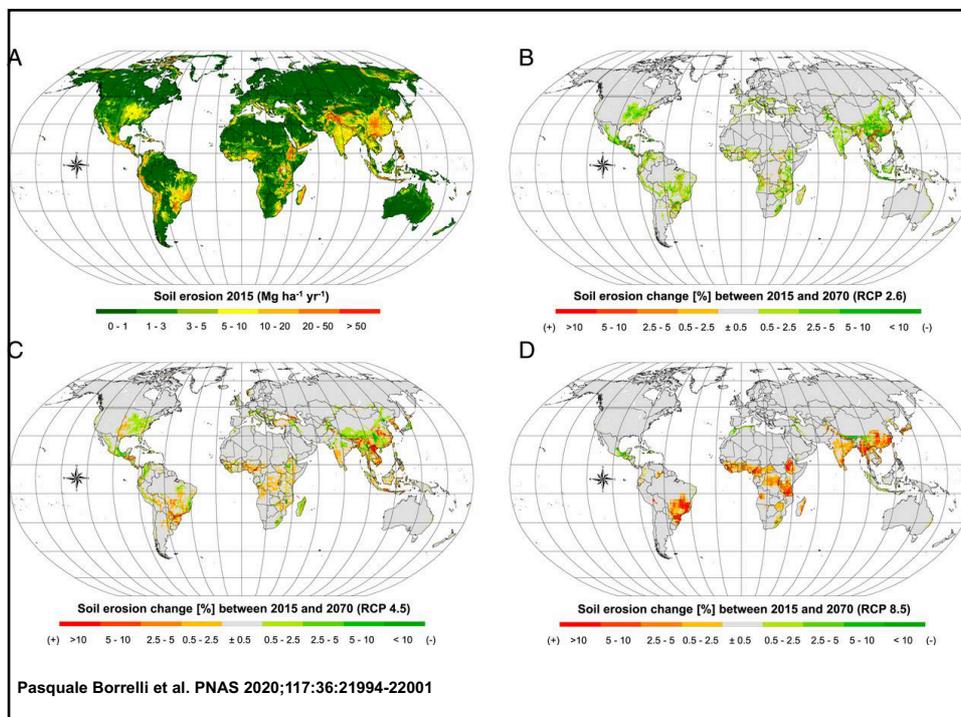
% Cumulativa em 25 anos

1	25
2	41
3	52
4	60
5	66

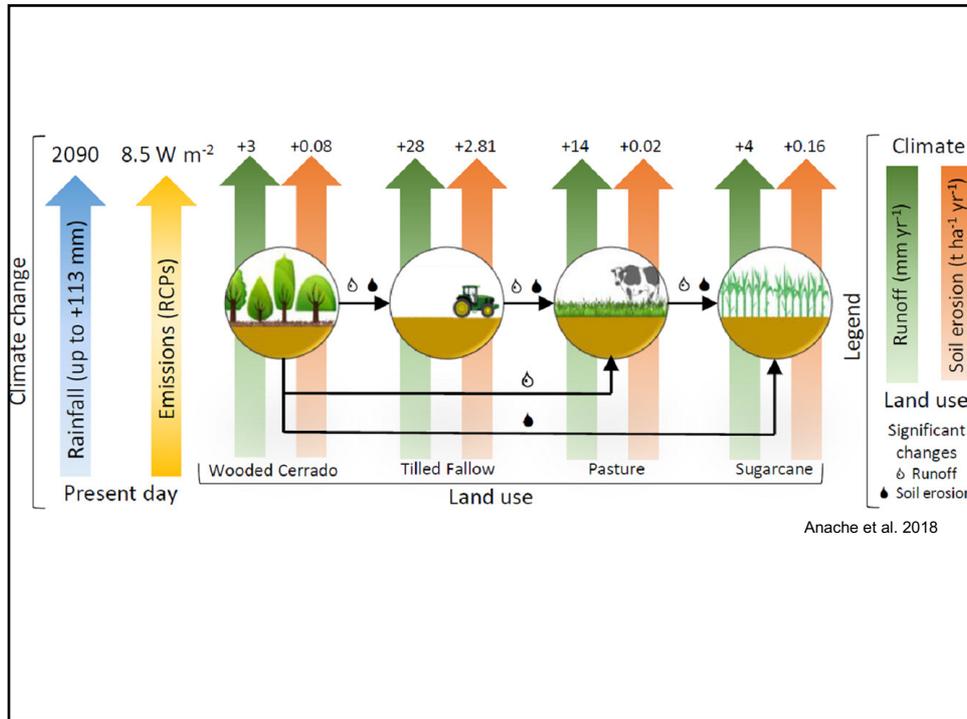
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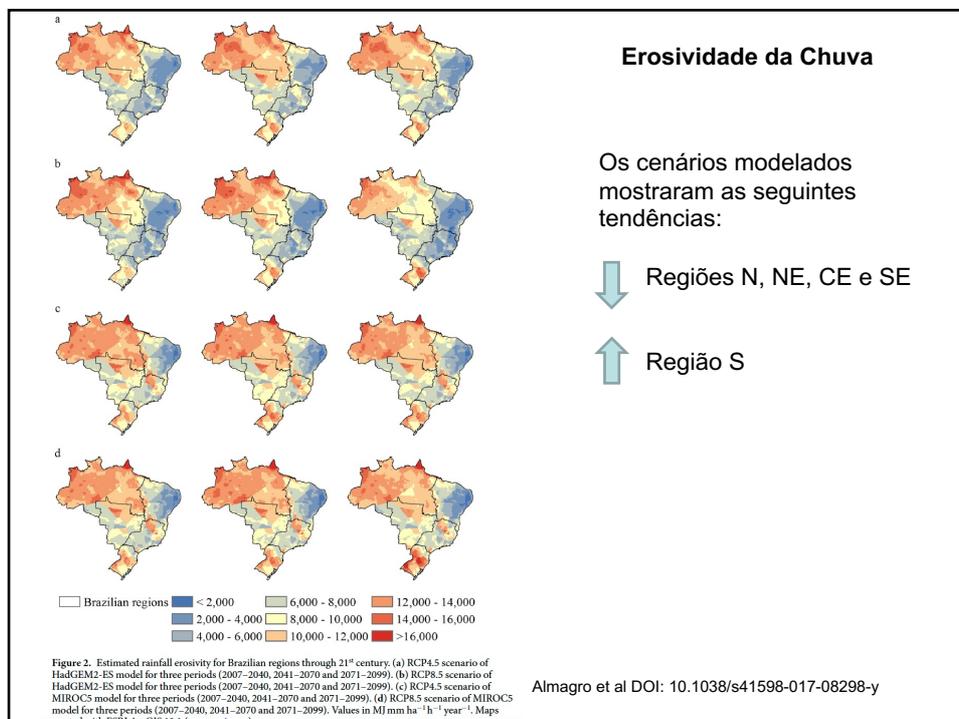
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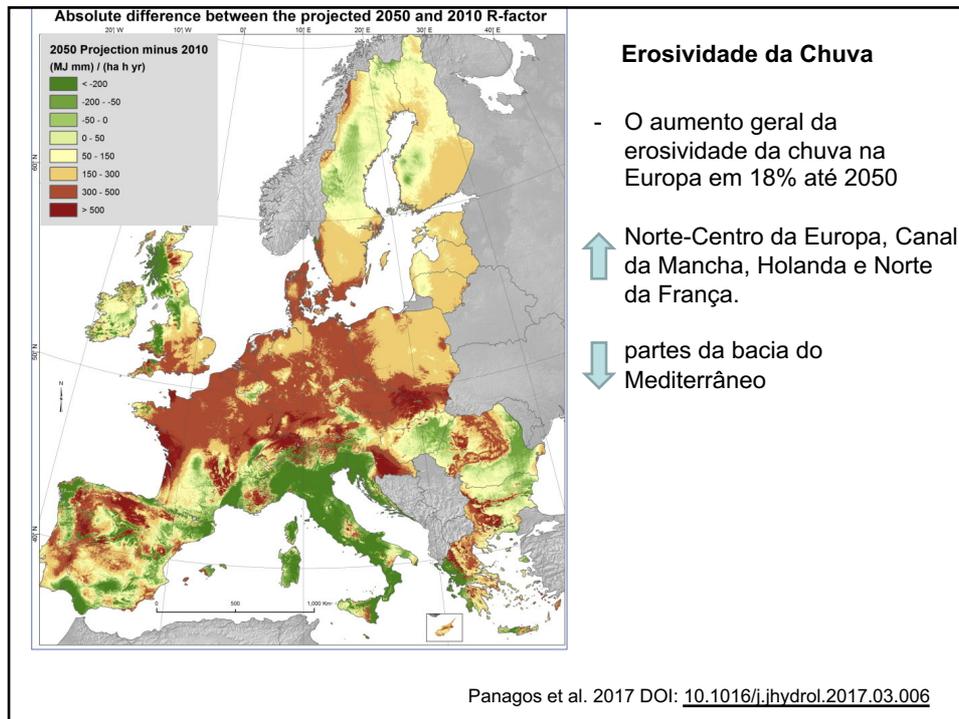
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